



JGR Biogeosciences

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RESEARCH ARTICLE

10.1029/2020JG005793

Key Points:

- Temperature, daylength, and light intensity influenced methane production from marine algae
- Light intensity showed a very pronounced effect on methane formation rates from all investigated species
- Global change will likely increase methane release rates of marine phytoplankton

Supporting Information:

· Supporting Information S1

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Citation:

Klintzsch, T., Langer, G., Wieland, A., Geisinger, H., Lenhart, K., Nehrke, G., & Keppler, F. (2020). Effects of temperature and light on methane production of widespread marine phytoplankton. *Journal of Geophysical Research: Biogeosciences*, *125*, e2020JG005793. https://doi.org/10.1029/2020JG005793

Received 22 APR 2020 Accepted 6 AUG 2020 Accepted article online 25 AUG 2020

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Effects of Temperature and Light on Methane Production of Widespread Marine Phytoplankton

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Abstract Methane (CH₄) production in the ocean surface mixed layer is a widespread but still largely unexplained phenomenon. In this context marine algae have recently been described as a possible source of CH₄ in surface waters. In the present study we investigated the effects of temperature and light intensity (including daylength) on CH₄ formation from three widespread marine algal species Emiliania huxleyi, Phaeocystis globosa, and Chrysochromulina sp. Rates of E. huxleyi increased by 210% when temperature increased in a range from 10°C to 21.5°C, while a further increase in temperature (up to 23.8°C) showed reduction of CH₄ production rates. Our results clearly showed that CH₄ formation of E. huxleyi is controlled by light: When light intensity increased from 30 to 2,670 μmol m⁻² s⁻¹, CH₄ emission rates increased continuously by almost 1 order of magnitude and was more than 1 order of magnitude higher when the daylength (light period) was extended from 6/18 hr light-dark cycle to continuous light. Furthermore, light intensity is also an important factor controlling CH₄ emissions of Chrysochromulina sp. and P. globosa and could therefore be a species-independent regulator of phytoplankton CH₄ production. Based on our results, we might conclude that extensive blooms of E. huxleyi could act as a main regional source of CH₄ in surface water, since blooming of E. huxleyi is related to the seasonal increase in both light and temperature, which also stimulate CH₄ production. Under typical global change scenarios, E. huxleyi will increase its CH₄ production in the future.

Plain Language Summary Methane is a gas that affects the Earth's climate and is typically produced by microbes in the absence of oxygen or through geological processes. Surprisingly, methane is also produced in oceanic surface waters that are well oxygenated, known as the ocean-methane paradox. Marine phytoplankton has recently been discovered as a methane source, which might help to explain the paradox. Environmental factors such as light and temperature might be important for controlling methane production from marine algae. In order to understand how environmental factors affect methane formation from phytoplankton, we performed several experiments under laboratory conditions. We find that temperature, light intensity, and day length strongly control methane production of phytoplankton. The field blooms of marine algae, which are often strongly related to the seasonal increase of light and temperature, could act as an important regional source of methane in oceanic surface waters. Under typical global change scenarios, marine algae might increase their methane production in the 21th century.

1. Introduction

Huge amounts of methane (CH_4) are formed in the oceans, but only a small proportion is released to the atmosphere (Weber et al., 2019). In this context the biogeochemical cycle of CH_4 in the oceans is of great interest, and in particular, the frequently observed CH_4 production within the ocean surface mixed layer is challenging our previous understanding of biogeochemical CH_4 formation processes. Traditionally, it is thought that CH_4 in the oceans is either produced by geological processes (abiotic) or by methanogenic archaea (biotic). Because methanogenic archaea are strict anaerobic microorganism, their CH_4 production is limited to anoxic environments (Kirschke et al., 2013; Saunois et al., 2016; Thauer et al., 2008). However, there is growing evidence that CH_4 is also produced by organisms such as cyanobacteria (Bižić et al., 2020) and eukaryotes including plants (Keppler et al., 2006), fungi (Lenhart et al., 2012), lichens



(Lenhart et al., 2015) and algae (Klintzsch et al., 2019; Lenhart et al., 2016), animals (Ghyczy et al., 2008), and humans (Keppler et al., 2016) and even in the presence of oxygen.

The observation of CH_4 in freshwater and saline surface waters (often described as methane paradox) has recently received much attention although some studies already conducted four decades ago (Scranton, 1977; Scranton & Brewer, 1977; Scranton & Farrington, 1977) have reported about CH₄ supersaturation in the ocean mixed layer. Furthermore, many recent studies (Grossart et al., 2011; Günthel et al., 2019; Hartmann et al., 2020; Tang et al., 2016) have shown that CH₄ formation is not limited to saltwater but also occurs in freshwater lakes. Several hypotheses exist to explain CH₄ formation in oxygenated waters, and some of them will be discussed briefly. Methanogenic archaea living in anoxic environments of particles or fish and zooplankton guts might form CH₄ (de Angelis & Lee, 1994; Karl & Tilbrook, 1994; Schmale et al., 2018; Stawiarski et al., 2019; Zindler et al., 2013). The algal methabolit dimethylsulfoniopropionate (DMSP) and its degradation products dimethyl sulfide (DMS) or dimethyl sulfoxide (DMSO) could be precursors of both archaeal (Damm et al., 2008; Florez-Leiva et al., 2013) and bacterial produced CH₄, when bacteria suffer under nitrogen deficiency (Damm et al., 2010). Moreover, photochemical degradation of DMS and acetone has been shown to produce CH₄, but the reaction is limited to anoxic waters (Bange & Uher, 2005; Zhang, Xie, et al., 2015). In oligotrophic Pacific waters CH₄ formation might mainly related to the bacterial cleavage of methylphosphonates when supply of phosphorous is limited (del Valle & Karl, 2014; Karl et al., 2008; Metcalf et al., 2012; Repeta et al., 2016).

Phytoplankton might contribute to CH₄ production in both oxic marine and freshwater environments. The first indication of CH₄ production from phytoplankton was provided by culture experiments of the diatom species Thalassiosira pseudonana and the haptophyte species E. huxleyi (Scranton, 1977; Scranton & Brewer, 1977; Scranton & Farrington, 1977). Later on, many field studies have reported a relationship between CH₄ supersaturation and the occurrence of phytoplankton in lakes and oceans (e.g., Bogard et al., 2014; Conrad & Seiler, 1988; Damm et al., 2008; Grossart et al., 2011; Oudot et al., 2002; Owens et al., 1991; Rakowski et al., 2015; Tang et al., 2014; Weller et al., 2013; Zindler et al., 2013). Although a good statistical correlation was not observed in all previous studies (e.g., Brooks et al., 1981; Burke et al., 1983; Forster et al., 2009; Lamontagne et al., 1975; Watanabe et al., 1995), it was suggested that phytoplankton is one of the likely CH₄ sources. However, clear evidence of CH₄ formation from marine algae—examined in cultures of marine haptophytes—was only provided recently when Lenhart et al. (2016) and Klintzsch et al. (2019) applied stable isotope labeling experiments to unambiguously show that the three widespread marine algae such as E. huxleyi, Chrysochromulina sp., and Phaeocystis globosa indeed produce CH4 per se and without the help of methanogenic archaea. Very recently, when using stable isotope labeling experiments and concentration measurements, it could be also shown that several freshwater algal species, including diatoms, cryptophytes, and green algae (Hartmann et al., 2020), but also several species of marine and limnic cyanobacteria (Bižić et al., 2020) emit CH₄. Thus, both marine algae and cyanobacteria could significantly contribute to the commonly observed oceanic CH₄ supersaturation (Bižić et al., 2020; Klintzsch et al., 2019; Scranton, 1977). In summary, previous investigations mainly focused on explaining the sources for CH₄ in oxic surface waters; however, the effects of environmental parameters such as temperature, light intensity, or nutrient availability on CH₄ production from phytoplankton are still unknown.

In the present study we investigated the effects of temperature and light intensity (including daylength) on CH_4 formation from the three widespread marine algal species *E. huxleyi*, *P. globosa*, and *Chrysochromulina* sp. *Emiliania huxleyi* occurs in ocean worldwide except in the polar regions (McIntyre et al., 1970). The algal species develops large populations (blooms) in subpolar to temperate areas usually in summer time, especially under highly stratified conditions, when the mixed layer depth shallows due to increasing temperature. Blooming of *E. huxleyi* is then supported by high light intensity caused by shallow mixed layer depth and incidence light (Nanninga & Tyrrell, 1996; Raitsos et al., 2006; Tyrrell & Merico, 2004; Tyrrell & Taylor, 1996). Therefore, we have studied in detail CH_4 formation in relation to temperature (range from $10.1^{\circ}C$ to $23.8^{\circ}C$), light intensity (30 to $2,670~\mu$ mol m⁻² s⁻¹), and daylength (period of light irradiation) during growth of *E. huxleyi*. We furthermore investigated the effect of light intensity on CH_4 formation by the two other widespread marine, but noncalcifying haptophytes *P. globosa* and *Chrysochromulina* sp. These two species can also form large blooms and are often found as main members in marine phytoplankton communities (Brown & Yoder, 1994; Schoemann et al., 2005; Thomsen, 1994). The results of *E. huxleyi* will be

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discussed with regard to their potential importance in marine environments during blooming. Finally, the observed CH_4 formation patterns of the three algal species will be evaluated on the basis of the CH_4 production potential (CH_4 -PP), which expresses differences in growth rates and thus the success of a species at the community level.

2. Materials and Methods

2.1. Experimental Setup

Emiliania huxleyi RCC1216 provided from the Roscoff Culture Collection (http://roscoff-culture-collection. org/, last access: 11 April 2020) were used to investigate the effect of temperature, light intensity, and daylength on CH₄ production rates. We performed an additional experiment to study the effect of light intensity on different algae species. Therefore, E. huxleyi and two other haptophytes P. globosa PLY 575 and Chrysochromulina sp. PLY 307 obtained from the Marine Biological Association of the United Kingdom (https://www.mba.ac.uk/facilities/culture-collection, last access: 11 April 2020) were studied. The cultures were maintained in quasi-exponential growth by frequent dilution with medium in order to keep them largely free of bacteria. All culture experiments were conducted under the use of sterile techniques. For a more detailed discussion about the potential interplay between algae and bacteria, we would like to refer the reader to the manuscript by Klintzsch et al. (2019). Briefly, Klintzsch et al. concluded that CH₄ production is clearly dependent on algal growth and that it is highly unlikely that bacteria alone are responsible for CH₄ production in the studied cultures. Each sample was taken at the end of the light period. Cultures were grown in batch mode (Langer et al., 2013). We used F/2 growth medium (Guillard & Ryther, 1962) that was based on sterile filtered (0.2 µm Ø pore size) North Sea seawater (sampled off Helgoland, Germany, 32 PSU). Cells were grown in crimped serum bottles (160 ml) filled with 140 ml medium and 20 ml headspace. Culture experiments were carried out with four independent repetitions. For determination of the CH₄ mixing ratio samples of 10 ml of headspace gas was sampled. The amount of produced CH₄ in culture group vials was calculated in respect to control groups. The culture and control group flasks were simultaneously sealed under ambient air and thus contained the same CH₄ background concentration. Please note that the produced CH₄ has been determined for the entire incubation flask—dissolved in the medium plus CH₄ of the headspace volume. For details on determination of CH₄ formation, please refer to section 2.2. For all experiments performed with algae and F/2 medium, the average CH₄ content in the cultures group at the end of incubation was higher than that found in the algae-free blanks as shown in Tables S1-S4. All growth rates and initial and final cell densities are given in the supporting information (Tables S1-S4).

Cultures were illuminated by cold white LED bulbs (LED Base Classic A100, Osram, Germany). The light spectrum of the LED bulbs is provided in Figure S1. The photosynthetic active radiation (PAR) was measured inside each incubation jar by using a light meter (ULM-500 Universal, WALZ, Germany) with a spherical quantum PAR sensor (US-SQS/L, WALZ, Germany). Temperature was logged by (UX120-006 M, HOBO, Germany).

2.2. Determination of CH₄ Mass

The CH₄ mass was determined at the end of the incubation period. In order to determine the CH₄ mass of the whole incubation flask (dissolved plus the CH₄ of the headspace volume), an aliquot (10 ml) of head space gas was taken from the incubation vials using a gastight syringe. In order to maintain headspace pressure when taking the headspace gas sample, an equivalent volume of seawater was injected into the flasks by syringe. The added volume was taken into account when determining the cell density (section 2.5). The sample gas was separated by gas chromatography using a GC-14B (Shimadzu, Japan) equipped with a 2 m column ($\emptyset = 3.175$ mm inner diameter), packed with Molecular Sieve 5A 60/80 mesh from Supelco. Methane was recorded by a flame ionization detector (FID) and quantified (mixing ratio) by using two reference standards containing 9,837 and 2,192 parts per billion by volume (p.p.b.v) CH₄. Mixing ratios were corrected for head space pressure. The latter was measured inside the incubation flask before gas sampling using a pressure meter (GMSD 1,3 BA, Greisinger). The CH₄ mass (m_{CH4}) was determined by its mixing ratio (x_{CH4}) and the ideal gas law (Equation 1),

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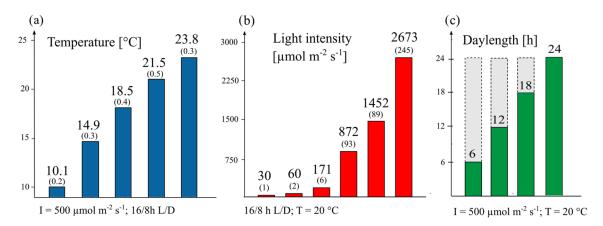


Figure 1. Treatment conditions of temperature (a), light intensity (b), and daylength (c) experiment. All treatments were carried out with four independent replications. L/D = light/dark; I = light intensity. (a) Mean values of logged temperature during the test period. (b) Mean light intensity values that were measured inside each incubation jar. The standard deviation is given in brackets.

$$m_{CH_4} = M_{CH_4} \times x_{CH_4} \frac{p \times V}{R \times T},\tag{1}$$

where M_{CH_4} = molar mass, p = pressure, T = temperature, R = ideal gas constant, and V = volume.

The dissolved CH₄ concentration was calculated by using the equation of Wiesenburg and Guinasso (1979).

2.3. Treatments of Alternating Temperature, Light Intensity, or Daylength at Cultures of *E. huxleyi*

To investigate the effect of temperature, light intensity, and daylength (daylength refers to the light period within a 24-hr light-dark cycle), three independent experiments with *E. huxleyi* were carried out in which one of the three parameters was varied as described in Figure 1. Within each experiment, the other two parameters were kept constant with 16/8-hr light-dark cycle; $\sim 500 \, \mu$ mol photons m⁻² s⁻¹ and 20° C, respectively. All treatments were carried out with four independent repetitions. Control groups contained F/2 medium only. Cultures were acclimated (≈ 10 generations) to the environmental conditions prior to the experiment. Cell density at inoculation varied between treatments depending on the growth rates under the given environmental conditions. Cultures of *E. huxleyi* were allowed to grow not more than 0.4×10^6 cells ml⁻¹ (exponential phase) before they were harvested. The majority (>95%) of culture replicates reached final cell densities between 0.1×10^6 and 0.3×10^6 cells ml⁻¹ (Tables S1–S3). Possible culture artifacts of CH₄ production rates, which could result from a cell density effect (Langer et al., 2013), were excluded for each investigated parameter by correlating the CH₄ production rates with the cell density on the harvest day (Figure S2).

2.4. Treatments of Alternating, Light Intensity on Cultures of *P. globosa, Chrysochromulina* Sp., and *E. huxleyi*

The effect of light intensity was studied in two further haptophytes: P. globosa and Chrysochromulina sp. in addition to E. huxleyi. Two light intensities $(427 \pm 12 \, \mu \text{mol m}^{-2} \, \text{s}^{-1})$ and $1,165 \pm 42 \, \mu \text{mol m}^{-2} \, \text{s}^{-1})$ with four replicates were applied, respectively. Cultures were grown under a 16/8 hr light-dark cycle and 20° C. Cultures were preadapted (≈ 10 generations) to light intensities before the experiment was started. The initial and final cell densities correspond to the exponential phase for each species (Klintzsch et al., 2019) and are given in the supporting information (Table S4). Control groups contained F/2 medium only.

2.5. Determination of Cell Density

For the determination of cell densities either a Fuschs-Rosenthal or Neubauer counting chamber (depending on cell density) was used. At least minimum of four aliquots of each culture sample were counted.

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2.6. Determination of Growth and CH₄ Production Rates

All production rates were measured at exponentially growing cultures. For further information of measuring production rates from batch culture experiments, we refer to Klintzsch et al. (2019) and Langer et al. (2012, 2013).

We calculated the growth rate (μ) from cell densities (N) of the beginning (t_0, N_0) and end (t_1, N_1) of the experiment (Equation 2).

$$\mu = \frac{Ln(N_1) - Ln(N_0)}{(t_1 - t_0)} \tag{2}$$

The POC-based CH_4 production rates were calculated from the cellular organic carbon content (POC_{cell}). The latter was obtained from cell volume (V_{cell}) by using the carbon to volume relationship in Equation 3 according to Menden-Deuer and Lessard (2000):

$$POC_{cell} = 0.216 \times V_{Cell}^{0.939}$$
 (3)

The cell volume was calculated from the cell diameter in light micrographs, which was measured by using the program ImageJ (Schindelin et al., 2012). We followed the recommendation of Olenina et al. (2006) and assumed a ball shape for calculating the cell volume for the three species investigated here

The carbon-specific growth rate was calculated from the product of POC and growth rate μ (Equation 4):

$$\mu_{POC} = \mu \times POC_{Cell}. \tag{4}$$

The CH₄ production rates were calculated by multiplying the growth rate μ with the corresponding cellular or POC-CH₄ quota, which was measured at the end of the experiment. The daily cellular CH₄ production rates (CH_4P_{cell} , ag CH₄ cell⁻¹ day⁻¹, ag = 10^{-18} g) were calculated according to Equation 5:

$$CH_4P_{cell} = \mu \times \frac{m(CH_4)}{cell},$$
 (5)

where $m(CH_4)$ is the amount of CH_4 that was produced at the end of the experiment.

The daily cellular CH₄ production rates (CH_4P_{POC} , μ g CH_4 g⁻¹ POC day⁻¹) were calculated from growth rate and CH₄-POC quotas at the end of the experiment according to Equation 6.

$$CH_4P_{POC} = \mu \times \frac{m(CH_4)}{POC}.$$
 (6)

The CH_4 production potential (CH_4 -PP) was calculated to scale variations in cellular production rates to community level. Detailed explanations for calculating the production potential (PP, which is not confined to CH_4) have been provided by Gafar et al. (2018) and Gafar and Schulz (2018). Please note that these authors have calculated the PP for $CaCO_3$ but the concept is the same for CH_4 . In accordance to the authors, the CH_4 -PP can be calculated for different growth periods, when a cellular standing stock for each time period is calculated from a given starting cell density (N_0). The related amount of produced CH_4 (CH_4 -PP) for each period of growth and respectively standing stock is the product of the cellular standing stock and CH_4 quota (Equation 7).

$$CH_4PP = N_0 \times e^{\mu \times t} \times \frac{m(CH_4)}{cell}$$
(7)

In the present study and in accordance to Klintzsch et al. (2019) the CH_4 -PP was calculated for a standing stock that is obtained after 7 days of growth starting with a single cell.

The sensitivity of growth, POC production and the rate of CH_4 formation to temperature were quantified by their activation energy (E_a), which is derived from the Arrhenius equation (Equation 8).

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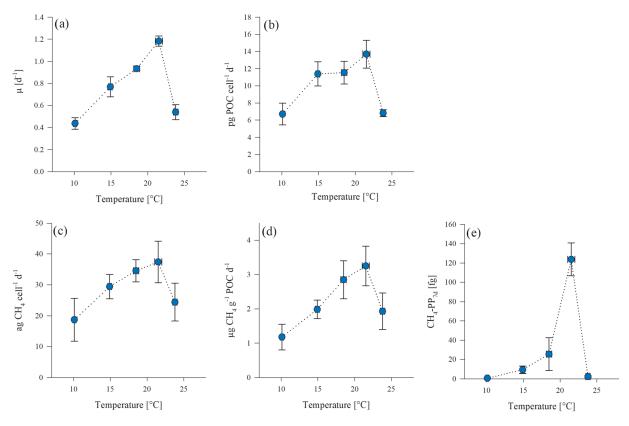


Figure 2. Relationship between temperature and growth rate (a), POC production rate (b), cellular (c), and POC normalized CH₄ production rate (d) and CH₄-PP (e) of *E. huxleyi*. Values are the mean of four replicated culture experiments with standard deviation (SD).

$$k(T) = A_{exp} \left(\frac{-E_a}{RT} \right), \tag{8}$$

where k = reaction rate constant (here for growth, POC, or CH₄ production rate), A = pre-exponential factor, R = gas constant, and T = temperature. The activation energies (E_a) of the rate can then be calculated by multiply the slope of the Arrhenius plot by -R, using a plot of $\ln(k)$ as function of T^{-1} .

2.7. Statistics

For each environmental factor (sections 3.1–3.3) the total data set of cellular and POC normalized CH_4 production was analyzed for statistical differences in the mean values among the treatment groups by using a one-way analysis of variance (ANOVA). Furthermore, within the individual culture experiments of *E. hux-leyi*, *P. globosa*, and *Chrysochromulina* sp. (section 3.4), the mean values (cellular and POC normalized CH_4 production) of the two light intensities treatments (medium and high light) were compared by t tests.

3. Results

3.1. Temperature Effect

Growth and POC production rates have more than doubled when increasing temperatures from 10.5° C to 21.5° C (Figures 2a and 2b). The optimum of growth and POC production was reached at 21.5° C ($1.18 \, day^{-1}$; $13.7 \pm 1.6 \, pg \, POC \, cell^{-1} \, day^{-1}$), while a further increase in temperature to 23.8° C led to a drastic reduction of about 50% for both growth and POC production rates. A similar pattern was observed for CH₄ production rates (cellular and POC normalized) as shown in Figures 2c and 2d. The POC normalized and cellular CH₄ production increased by 2.8- and 2.0-fold, respectively, when temperature increased by 11.4° C (from 10.1° C to 21.5° C). At 21.5° C the optimum of CH₄ production rates was reached ($3.2 \pm 0.6 \, \mu g \, CH_4 \, g^{-1} \, POC \, day^{-1}$; $37.4 \pm 6.7 \, ag \, CH_4 \, cell^{-1} \, day^{-1}$). Further increase in temperature from 21.5° C to 23.8° C

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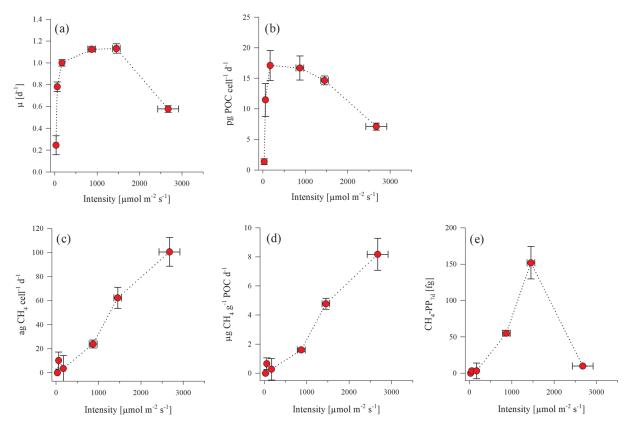


Figure 3. Relationship between light intensity and growth rate (a), POC production rate (b), cellular (c), and POC normalized CH₄ production rate (d) and CH₄-PP e of *E. huxleyi*. Values are the mean of four replicated culture experiments with SD.

showed a reduction of $\mathrm{CH_4}$ production by 40% and 35%, for POC normalized and cellular $\mathrm{CH_4}$ production, respectively. Statistical analysis (ANOVA) confirmed the temperature dependence of $\mathrm{CH_4}$ production with p values of 0.002 and <0.001 for POC normalized and cellular $\mathrm{CH_4}$ production rates, respectively. After 1 week of growth the total amount of generated $\mathrm{CH_4}$ is specified by the $\mathrm{CH_4}$ -PP (Figure 2e). With increasing temperature (from 10.1°C to 21.5°C) the $\mathrm{CH_4}$ -PP raised by 2 orders of magnitude (from 0.7 \pm 0.2 to 124 \pm 17.1 fg $\mathrm{CH_4}$) before it declined drastically at 23.8°C (Figure 2e). Consequently, the optimum temperature (21.5°C) was identical for the five investigated parameters (Figure 2).

3.2. Light Intensity Effect

The growth and POC production rates of *E. huxleyi* increased drastically when light intensity increased from 30 to 171 μ mol m⁻² s⁻¹, and values remained relatively constant at higher light intensities in the range of 171–1,450 μ mol m⁻² s⁻¹ (Figures 3a and 3b). However, increasing the light intensity to 2,670 μ mol m⁻² s⁻¹ caused a clear reduction of both growth and POC production rates. From 30 to 171 μ mol m⁻² s⁻¹ the growth rate increased fourfold (from 0.25 \pm 0.09 to 1.00 \pm 0.03 day⁻¹) and POC production rates by over 1 order of magnitude (from 1.4 \pm 0.48 to 17.1 \pm 2.5 pg POC cell⁻¹ day⁻¹). The POC production was highest at 171 μ mol m⁻² s⁻¹ and was similar at 872 μ mol m⁻² s⁻¹ while rates were slightly smaller at 1452 μ mol m⁻² s⁻¹. The growth rate increased by 10% between 171 and 872 μ mol m⁻² s⁻¹ but did not further change when reaching values of 1,450 μ mol m⁻² s⁻¹ (1.13 \pm 0.03 day⁻¹). Thus, the optimum growth rate is reached at higher light intensities compared to the POC production rates. However, a further increase in light intensity up to 2,673 μ mol m⁻² s⁻¹ led to a significant reduction (\approx 45%) of both POC production and growth rates. In contrast, the cellular and POC normalized CH₄ production rates increased steadily with increasing light intensity (Figures 3c and 3d). Methane formation (on a cell basis) was below the detection limit at 30 μ mol m⁻² s⁻¹ but measurable (6.9 \pm 9.5 ag CH₄ cell⁻¹ day⁻¹; 0.47 \pm 0.63 pg POC cell⁻¹ day⁻¹ on average) between 60 and 171 μ mol m⁻² s⁻¹ while at 872 μ mol m⁻² s⁻¹ production rates strongly increased (23.9 \pm 3.3 ag CH₄ cell⁻¹

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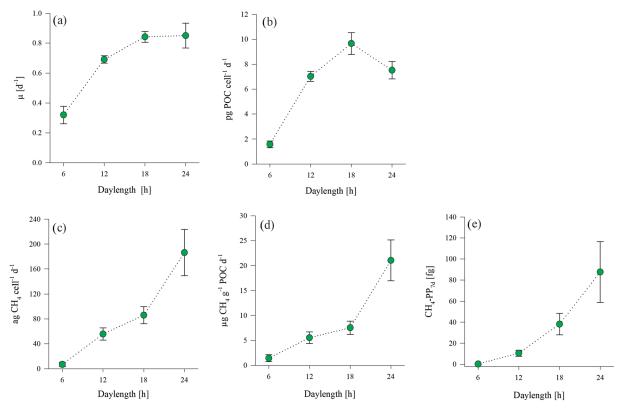


Figure 4. Relationship between day length and growth rate (a), POC production rate (b), cellular (c), and POC normalized CH₄ production rate (d) and CH₄-PP (e) of *E. huxleyi*. Values are the mean of four replicated culture experiments with SD.

day $^{-1}$; 1.6 \pm 0.1 μ g CH $_4$ g $^{-1}$ POC day $^{-1}$). From 872 to 2,670 μ mol m $^{-2}$ s $^{-1}$ cellular and POC normalized CH $_4$ production increased by 4.2- and 5.1-fold up to 100 \pm 12 ag CH $_4$ cell $^{-1}$ day $^{-1}$ and 6.8 \pm 0.9 μ g CH $_4$ g $^{-1}$ POC day $^{-1}$, respectively. The light dependence of both cellular and POC-normalized CH $_4$ production was also indicated by statistical analysis (ANOVA; p < 0.001). The CH $_4$ -PP (Figure 3e) increased with increasing light intensities by 2 orders of magnitude up to 152 \pm 22 fg CH $_4$ at 1450 μ mol m $^{-2}$ s $^{-1}$ and sharply decreased by on order of magnitude (to 9.9 \pm 0.9 fg CH $_4$) at higher light intensity (2,670 μ mol m $^{-2}$ s $^{-1}$). The optimum of CH $_4$ -PP is therefore in accordance with the optimum of growth rate.

3.3. Daylength Effect

The extension of the daylength (period of light irradiation) from 6 to 18 hr increased the growth rates 2.6-fold (from 0.32 ± 0.06 to 0.84 ± 0.04 day⁻¹), while the growth rates remained constant when a period of continuous light (24 hr) was set (Figure 4a). In contrast to the growth rate, an optimum of POC production rates was observed at 18 hr daylength and decreasing with longer irradiation period of 24 hr (Figure 4b). POC production rates increased between 6 and 18 hr daylength by 6.1-fold (from 1.6 ± 0.3 to 9.7 ± 0.9 pg POC cell⁻¹ day⁻¹) and declined by 22% at continuous light. Cellular and POC normalized CH₄ production rates increased from 6 hr daylength to continuous light period by 2 and 1 order of magnitude from 6.9 ± 3.4 to 186 ± 37 ag CH₄ cell⁻¹ day⁻¹ and 1.1 ± 0.6 to 21.0 ± 4.1 µg CH₄ g POC⁻¹ day⁻¹, respectively (Figures 4c and 4d). The dependence of CH₄ production on temperature was verified by statistical analysis (ANOVA), with cellular and POC normalized CH₄ production rates showing *p* values of *p* < 0.001 and *p* = 0.004, respectively. The cellular and POC normalized CH₄ production was particularly enhanced by the 6 hr extension of the daylength between 18 hr and continuous light that accounted for 56% and 69% of the total increase in cellular and POC normalized CH₄ production, correspondingly. The CH₄-PP increased constantly by over 2 orders of magnitude with longer light irradiation periods (from 0.19 \pm 0.12 fg CH₄ at 6 hr light to 70.4 \pm 23.1 fg CH₄ at continuous light, Figure 4e).

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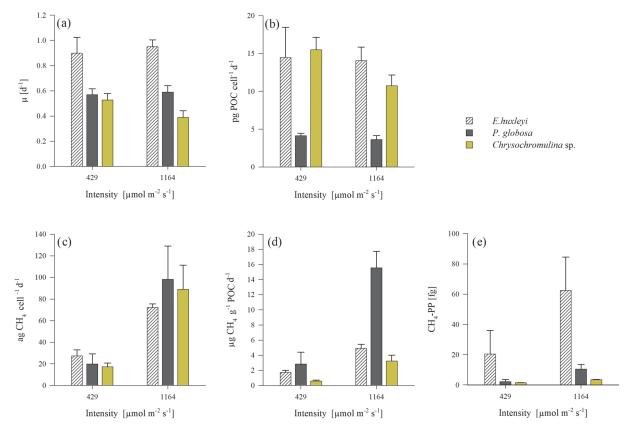


Figure 5. Relationship between light intensity and growth rate (a), POC production rate (b), cellular (c), and POC normalized CH₄ production rate (d) and CH₄-PP (e) of *E. huxleyi*, *P. globosa*, and *Chrysochromulina* sp. by moderate (429 μ mol m⁻² s⁻¹) and high light intensity (1,164 μ mol m⁻² s⁻¹). Values are the mean of four replicated culture experiments with SD.

3.4. Comparison of Light Intensity Effects of E. huxleyi, Chrysochromulin Sp., and P. globosa

We compared growth and CH₄ formation patterns of the three algal species E. huxleyi, P. globosa, and Chrysochromulina sp. at moderate and high light intensities (429 μmol m⁻² s⁻¹ and 1,164 μmol m⁻² s⁻¹). The growth rates at both light intensities are shown in Figure 5a. At moderate light intensity the exponential growth rate μ was highest for *E. huxleyi* (0.90 \pm 0.13 day⁻¹) followed by *P. globosa* and *Chrysochromulina* sp. (with $0.57 \pm 0.05 \text{ day}^{-1}$ and $0.55 \pm 0.04 \text{ day}^{-1}$, respectively). Growth rates of E. huxleyi and P. globosa remained constant at higher intensity, while growth rate of Chrysochromulina sp. declined by 29%. The POC production rates are shown in Figure 5b. At moderate light intensity the POC production rates of E. huxleyi and Chrysochromulina sp. were in the same range with 14.5 ± 0.5 and 15.5 ± 1.6 pg POC cell⁻¹ day^{-1} , respectively, and were about three times higher than for P. globosa (4.1 ± 0.3 pg POC cell⁻¹ day⁻¹). The exposure to high light intensity led to a 31% lower POC production rate of Chrysochromulina sp. while rates of E. huxleyi and P. globosa remained constant. Thus, an increase in light intensity declined growth rate and POC production of Chrysochromulina sp., while that of E. huxleyi and P. globosa remained constant. Cellular CH₄ production rates of all investigated species were enhanced by increasing light intensities (Figure 5c). Cellular CH₄ production rates ranged from 17 \pm 3.6 (Chrysochromulina sp.) to 27 \pm 5.6 ag $CH_4 cell^{-1} day^{-1}$ (E. huxleyi) at medium light. In response to higher light intensity the cellular CH_4 production rates increased by 2.6-fold (E. huxleyi) and about fivefold (P. globosa and Chrysochromulina sp.) resulting in a cellular CH₄ production rates ranged from 72.1 ± 3.3 (E. huxleyi) to 98.2 ± 30.8 ag CH₄ cell⁻¹ day⁻¹ (P. globosa). The response in cellular production rates to higher light intensity was also displayed by t test with p values of <0.001, 0.006, and 0.005 for E. huxleyi, P. globosa, and Chrysochromulina sp., respectively. The POC normalized CH₄ production rates increased with increasing light intensity (Figure 5d). Within the medium and high light intensities, the variation of POC normalized CH4 production rates between

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species was greater than that of cellular CH_4 production rates. When CH_4 production rates were normalized to POC, rates of moderate light intensities were in a range of 0.58 ± 0.12 to 2.8 ± 1.6 µg CH_4 g⁻¹ POC day⁻¹ with *Chrysochromulina* sp. and *P. globosa* showing the lowest and highest rates, respectively. At high light intensity rates were about 2.8-fold (*E. huxleyi*) and 5.5-fold (*P. globosa* and *Chrysochromulina* sp.) greater than for those observed at moderate light intensity. These differences were also shown by *t* test with *p* values <0.001, 0.004, and <0.001 for *E. huxleyi*, *P. globosa*, and *Chrysochromulina* sp., respectively. The respectively rates ranged from 3.24 \pm 0.78 (*Chrysochromulina* sp.) to 15.6 ± 2.2 µg CH_4 g⁻¹ POC day⁻¹ (*P. globosa*). All three species showed enhanced CH_4 -PP with the higher light intensity (Figure 5e). The increase from moderate to high light ranged between 2.4-fold (*Chrysochromulina* sp.) and 4.9-fold (*P. globosa*). However, the variation of the CH_4 -PP within the species is greater than that resulting from the different light treatments. The CH_4 -PP was 1 order of magnitude higher for *E. huxleyi* in comparison to the other two species (Figure 5e). This is in line with the higher growth rate of *E. huxleyi*.

4. Discussion

Previous studies indicated that several marine algae produce CH₄ (Klintzsch et al., 2019; Lenhart et al., 2016; Scranton, 1977; Scranton & Brewer, 1977; Scranton & Farrington, 1977), while the modulating influence of environmental parameters is unknown. Our results clearly show that CH₄ formation by *E. huxleyi* is influenced by temperature, light intensity, and the length of irradiation period. Furthermore, light intensity is also an important factor controlling emission rates of the two other marine algae *Chrysochromulina* sp. and *P. globosa*. We will first discuss the effects of environmental parameters on growth and POC normalized CH₄ production from a physiological perspective. Afterward, the effects of environmental parameters on laboratory CH₄ production rates of *E. huxleyi* are discussed in relation to their possible importance on populations (blooms) in marine environments. Finally, we discuss the impact of environmental parameters on CH₄ production in biogeochemical terms using the well-established but rarely applied concept of the PP (see Klintzsch et al., 2019, and references therein).

4.1. Temperature Effect on Growth and CH₄ Formation of *E. huxleyi* From a Physiological Perspective

Emiliania huxleyi occurs, except for the polar regions, in oceans worldwide and has the largest known temperature growth range (1–31°C) compared to other coccolithophores (McIntyre et al., 1970). The temperature response of growth rate is strain specific (Brand, 1982; Langer et al., 2009), and the optimum temperature for strain RCC1216 in this study tallies well with the published value (Langer et al., 2009). The growth curve (Figure 2a) exhibits the asymmetry typical for a temperature response. The ascending, shallow sloped, part of the curve is characterized by an accelerating effect of temperature on all biochemical reactions, whereas the descending, steep sloped, part is characterized by inactivation of enzymes, and denaturation of proteins and membranes (DeLong et al., 2017; Grimaud et al., 2017; Kingsolver, 2009). In accordance with this general concept of temperature effects on physiological processes, we observe a positive correlation of all analyzed physiological parameters with temperature up to 21.5°C (the optimum) followed by a negative correlation above this temperature. We conclude that CH₄ production is a normal physiological process as opposed to a heat stress response stemming from structural damage to cellular architecture. Please note that CH₄ production trends are identical, regardless of the normalization, that is, normalization to cell or POC (Figures 2c and 2d).

The ascending part of the temperature curve can be further analyzed using the Arrhenius equation. According to this equation (Equation 8), the thermal sensitivity of a chemical reaction is proportional to its activation energy. While the Arrhenius equation was originally used to describe chemical reactions, the equation might be also applied to describe the thermal sensitivity of biochemical reactions and biological growth rates, whereby high activation energies indicate high sensitivity to temperature (Gillooly et al., 2001; Grimaud et al., 2017). The calculated activation energies of growth rate, POC, and CH_4 production were 59, 41, and 63 kJ mol^{-1} , respectively. The growth rate and CH_4 production are therefore somewhat more sensitive to temperature than POC production is. The activation energy of CH_4 production is in the range of basic metabolic processes, indicating that CH_4 production in algae is not an abiotic process. For example, the average activation energy of respiration for a wide range of organisms, including microbes, plants, and animals,

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is between 40 and 71 kJ mol $^{-1}$ (Gillooly et al., 2001). In addition, activation energies of most enzymatic reactions are in the range of 21 to 63 kJ mol $^{-1}$ (Segel, 1993). By contrast abiotic CH₄ formation from thermal degradation experiments as described from dried soils usually showed higher activation energies above 70 kJ mol $^{-1}$ (Jugold et al., 2012; Liu et al., 2019).

4.2. Light Intensity and Daylength Effects on Growth and CH₄ Formation of *E. huxleyi* From a Physiological Perspective

We grew E. huxleyi under a wide range of light intensities and daylengths.

4.2.1. Light Intensity

Our results demonstrate that CH₄ formation and growth rate of E. huxleyi is sensitive to light intensity. Under light-limited condition, growth rate increased sharply with increasing light intensity, leveled off at saturated light, and decreased at inhibiting light intensities (Figures 3a and 3b). This pattern of light intensity response (Figure 3a) is typical of phytoplankton cultures (Edwards et al., 2015, and reference inside). The optimum light intensity for POC production is lower than the one for growth rate, which was also observed by Trimborn et al. (2007). The growth rate of E. huxleyi was remarkably tolerant against high light intensities ($\geq 1,500 \ \mu \text{mol m}^{-2} \ \text{s}^{-1}$), a phenomenon well documented in the literature (Balch et al., 1992; Gafar & Schulz, 2018; Harris et al., 2005; Loebl et al., 2010; Nanninga & Tyrrell, 1996; Nielsen, 1997; Trimborn et al., 2007). Interestingly CH₄ production was even more tolerant to high light, so much so that we could not determine the optimum light intensity. This is in notable contrast to the temperature response patterns described above. While we do not know the chain of events leading to this light intensity response, the response patterns suggest that, first, CH₄ production is a light dependent process and, second, photo-inhibition of growth rate and POC production do not impair CH₄ production. The latter is particularly intriguing because it seems to suggest a decoupling of CH₄ production from photosynthetic production of both energy equivalents and putative CH₄ precursors originating in the POC pool. This warrants further, more detailed, physiological studies into the nature of the light dependency of CH₄ production.

4.2.2. Daylength

The CH₄ formation and growth rate were furthermore controlled by daylength. The growth rate showed a saturation curve (Figure 4a). This pattern is similar to results on E. huxleyi reported by Paasche (1967) and even other phytoplankton species (e.g., Bouterfas et al., 2006). By contrast, growth rates of E. huxleyi have been reported to be independent of daylength (Nielsen, 1997) or to be inhibited by continuous light (Van Rijssel & Gieskes, 2002). The response to daylength has been suggested to be strain specific (Bretherton et al., 2019) and is dependent on other environmental parameters, for example, on seawater CO2 concentration (Bretherton et al., 2019; Zhang, Bach, et al., 2015) and light quality (Glover et al., 1987). This could be one reason why the response of growth in relation to daylength differs between studies. In our study POC production decreased at continuous light while growth rate did not (Figure 4b). While growth rate might be inhibited by a lack of a dark period rather than by photoinhibition (Brand & Guillard, 1981), the decline of POC production at continuous light could partly be due to photoinhibition. Please note (see also above) that the response pattern of CH₄ production is independent of the normalization (cell or POC). Interestingly, it is again CH₄ production that neither levels off nor shows inhibition at continuous light. This observation reinforces the pattern described above, namely, POC production declines while CH₄ production increases further. This "double dependency" of CH₄ production on light, that is, both light intensity and daylength, renders light-dependent processes the prime target for further elucidating the mechanism of CH₄ production in *E. huxleyi*.

4.3. Light Intensity Effects on Growth and CH_4 Formation of E. huxleyi, Chrysochromulin Sp., and P. globosa

Methane production was light dependent in P. globosa and Chrysochromulina sp. too. $Emiliania\ huxleyi$ and P. globosa showed a greater light tolerance with respect to growth rate and POC production than Chrysochromulina sp. (Figures 5a and 5b). In contrast to POC production and growth rate of Chrysochromulina sp., CH_4 production was not inhibited and once again confirms the remarkable light dependency of CH_4 production. The increase in CH_4 production with light intensity in P. globosa and Chrysochromulina sp. was even higher than that of E. huxleyi. However, in each case there was a positive correlation of CH_4 production and light intensity, which could therefore be a common feature of different

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phytoplankton taxa. This hypothesis is supported by recent findings of Bižić et al. (2020), who investigated different cyanobacterial species that are found in phytoplankton communities of ocean and lakes. While cyanobacteria produce CH_4 in light and dark phase, the CH_4 production rates elevated during the light phase.

4.4. Potential Relevance for CH₄ Production of E. huxleyi Populations in the Field

Large blooms of E. huxleyi typically occur at subpolar to temperate areas in the summer months, when the water is highly stratified, due to the seasonal increase in temperature and light intensity (Iglesias-Rodríguez et al., 2002; Nanninga & Tyrrell, 1996; Raitsos et al., 2006; Tyrrell & Merico, 2004). The occurrence of E. huxleyi in field is therefore correlated with high solar radiation, shallow mixed layer depth, and increased sea surface temperature (SST; Raitsos et al., 2006). We compare the reported environmental conditions that support E. huxleyi growth in the field with those that stimulate CH_4 production in our laboratory grown cultures to assess whether CH_4 formation by E. huxleyi could be of ecological relevance.

Emiliania huxleyi grows at water temperatures between 1°C and 31°C in field and has the largest temperature growth range of all coccolithophores (McIntyre et al., 1970). The wide temperature range results from the adaptation of individual strains to narrower temperature ranges in cold or warm water masses (Brand, 1982; Langer et al., 2009). The temperature range of the incubation experiments includes the range of the seasonal variation in the SST of the Tasman Sea off New Zealand where the investigated strain (RCC1216) has been isolated and is therefore of ecological relevance. The monthly mean SST of the Tasman Sea off New Zealand ranges from 13.5°C to 18.7°C between coldest and warmest month (time period 2007–2017, https://statisticsnz.shinyapps.io/sea_surface_temperature_oct19/, last access: 11 April 2020). With a temperature increase from 10.1°C to 21.5°C cellular CH₄ production rates would double according to laboratory derived rates (section 3.1). Inhibition of growth rate and CH₄ production due to heat stress is less likely in the field, since the mean SST of the warmest month is below the optimum temperature of the investigated strain. This is in line with literature data showing that most strains grow below their optimal growth condition in field (Langer et al., 2009; Rosas-Navarro et al., 2016).

Our laboratory data also suggest that longer light irradiation periods during summer could have a stimulating effect on CH₄ production, especially on E. huxleyi populations in subpolar regions, where daylength changes dramatically between winter and summer. For example, the CH₄ production would increase by a factor of 5 due to a daylength increase from 6 hr in winter to 18 hr in summer (section 3.3). Emiliania huxleyi usually blooms in the North Atlantic in June and July at high light intensity in the surface layer, which is caused by strong sunlight and shallow mixed layer depth (10-20 m; Nanninga & Tyrrell, 1996; Raitsos et al., 2006; Tyrrell & Merico, 2004; Tyrrell & Taylor, 1996). For instance, light intensities of 935 and 1,140 μ mol m⁻² s⁻¹ were measured in *E. huxleyi* blooms in the field. In addition, long-term observations in a mesocosm in a Norwegian fjord have shown that E. huxleyi blooms at light intensities between >530 and 1,200 µmol m⁻² s⁻¹, whereas the mean light intensities in the surface layer are ~63% and 43% of the incident light intensity at 10 and 20 m mixed layer depth, respectively (Nanninga & Tyrrell, 1996, and reference inside). Thus, the E. huxleyi cells would have been exposed to light intensities between 228 and 756 μ mol m⁻² s⁻¹ in the mixed layer, which falls within the light intensity range of our incubation experiments (section 3.2). Since the CH₄ production of E. huxleyi increased linearly with the light intensities in culture experiments, the high light intensities in the surface layer could also support the CH₄ formation of E. huxleyi in the field. Judging from our laboratory CH₄ production and the reported light intensity range where blooms typically occur, the CH₄ production could vary by a factor of 4. It can be concluded that light intensity will considerably affect sea surface water CH₄ production of E. huxleyi in field.

4.5. The Biogeochemical Perspective: Methane PP (CH₄-PP)

While the considerations in section 4.4 apply to the behavior of field populations on the cellular level, they are not appropriate for assessing the biogeochemical significance of this behavior. Several recent studies have emphasized that the PP (see section 2 for calculation), as opposed to the cellular production, is the relevant parameter for biogeochemical assessments (Gafar & Schulz, 2018; Gafar et al., 2018; Klintzsch et al., 2019; Kottmeier et al., 2016; Marra, 2002; Schlüter et al., 2014). We calculated the CH₄-PP of *E. huxleyi* for different temperature, light intensity, and daylength conditions. For all three parameters, the CH₄-PP increases toward the optimum, as does the cellular CH₄ production, but the increase in CH₄-PP was by 1 order of magnitude higher than for the cellular CH₄ production (Figures 2c, 3c, and 4c and Figures 2e, 3e,

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and 4e). This illustrates the importance of using the PP when considering the biogeochemical impact of changing environmental conditions. Another such illustration is the strong contrast between the light intensity response patterns of cellular, or POC normalized, CH₄ production, and the CH₄-PP. The sharp decline in CH₄-PP at the highest light intensity is not reflected in the cellular CH₄ production curve. However, in the field this difference is of minor importance because the highest light intensity used here, ~2,700 µmol m⁻² s⁻¹, is considerably higher than even peak light intensities observed in typical E. huxleyi blooms (~1,200 µmol m⁻² s⁻¹; see references above). It is noteworthy that the decline in CH₄-PP at the highest temperature tested here is also of little relevance in the field because E. huxleyi usually grows at suboptimal temperatures in the field, a situation that will also not change in the foreseeable future, despite global warming (Rosas-Navarro et al., 2016). It is concluded first that the CH₄-PP of E. huxleyi in the field will be maximal in midsummer when E. huxleyi typically blooms. Second, global change will increase the CH₄-PP of E. huxleyi through both warming and increased stratification entailing higher light intensities in the surface layer. Compared to the other two tested haptophytes, E. huxleyi has the highest CH₄-PP, a difference not mirrored in cellular CH₄ production: This is yet another example of the importance of using the PP when considering the biogeochemical impact of CH₄ formation by phytoplankton. As a general caveat it should be noted that the above conclusions are confined to our experimental conditions. Conditions in the field will include other factors such as grazing. This inevitable limitation of experimental data is the price one has to pay for discovering relationships between environmental parameters and the performance of an organism. However, this does not mean that our data are never directly applicable to the field situation as illustrated by the good match of satellite data and E. huxleyi calcite production potential reported by Gafar et al. (2018).

5. Conclusions

We have determined the CH₄ production of three haptophytes under varying environmental conditions and conclude the following:

- 1. Temperature, light intensity, and daylength influence CH₄ production.
- 2. CH₄ production is strongly light dependent; even increasing with light intensity when growth rate and POC production are photoinhibited.
- 3. The biogeochemically relevant parameter CH_a-PP increased with temperature, light intensity, and daylength over the range typical for present-day seasonality and global change predictions for the coming
- 4. E. huxleyi has a considerably higher CH₄-PP than P. globosa and Chrysochromulina sp.

Conflict of Interest

The authors declare that they have no conflict of interest.

Data Availability Statement

We provide the data in heiDATA, which is an institutional repository for research data of Heidelberg University (https://doi.org/10.11588/data/AGKWSG).

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Acknowledgments

We thank Dr. Markus Greule, Bernd Knape, and Stefan Rheinberger for conducting analytical measurements and for technical support that helped to produce this data set. Furthermore, we are grateful to Dr. Katja Grossmann and Dr. Stefan Schmitt for spectral measurements and Dr. Steffen Greiner for providing microscopy facilities. This work was supported by the Deutsche Forschungsgemeinschaft (Grant Nos. KE 884/11-1 and KE 884/16-2) and the Natural Environment Research Council (NE/N011708/1).

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